Crust breaking strength and durability

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Introduction: motivation

Today's key questions:

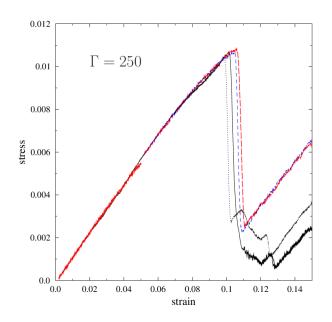
 What is the minimum, typical, and maximum ellipticity one should expect?
 What is the strength (breaking strain) of the crust?

 How does strain evolve in the crust and how does it break?

- What are the implications of observed upper limits on the ellipticity?

- At what point do upper limits become "interesting" (e.g. constrain theory)?

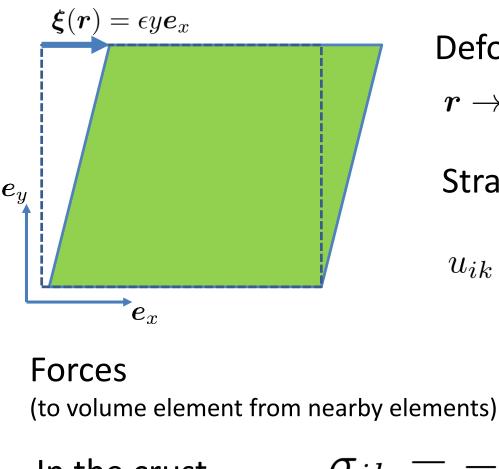
- What are possible mountain building mechanisms (such as asymmetric accretion, temperature gradients, magnetic stress...)?



Typical stress-strain curves [MD simulations]

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Introduction: Theory of Elasticity



Deformation:

$$m{r}
ightarrow m{r} + m{\xi}(m{r})$$

Strain tensor

$$u_{ik} = u_{ki} = \frac{1}{2} \left(\frac{\partial \xi_i}{\partial x_k} + \frac{\partial \xi_k}{\partial x_i} \right)$$

$$F_i = \frac{\partial \sigma_{ik}}{\partial x_k}$$

In the crust

$$\sigma_{ik} = -P\delta_{ik} + \Delta\sigma_{ik}$$

(electrons+neutrons+undeformed lattice)

 \sim Elastic part $\Delta \sigma_{ik}(u_{ik})$

Elastic part of stress tensor

Microphysics:

$$\Delta \sigma_{ik} = \Delta \sigma_{ik}(u_{ik}, \ldots)$$

$$\Delta \sigma_{ik} \lesssim 10^{-3} P_e$$

For infinitesimal deformations
$$\Delta \sigma_{ik} = \lambda_{iklm} u_{lm}$$

Elastic coefficients (up to 21)

For cubic crystal 3 independent coefficients (2 for isotropic material):

$$c_{11} = \lambda_{xxxx}, \quad c_{12} = \lambda_{xxyy}, \quad c_{44} = \lambda_{xyxy}$$

Basic model and scaling

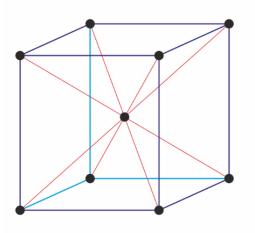
[one component crust – all ions are equal]

$$\Delta \sigma_{ik} = \frac{Z^2 e^2}{a} n \Delta \tilde{\sigma}_{ik} \left(u_{ik}, \Gamma = \frac{Z^2 e^2}{aT}, \frac{T}{\hbar \omega_P}, ak_{\rm TF}, \ldots \right)$$

Point charges, TF screening:

$$a = (4\pi n/3)^{-1/3}$$

Ions form BCC lattice



(Arbitrary) uniform
compression/expansion does
not lead to breaking
(but can initiate to nuclear reactions)

<u>Beyond scope:</u> Realistic (Jancovici 1962) electron screening [Baiko 2002] Effects of free neutrons: induced interactions [Kobyakov&Pethick 2016]

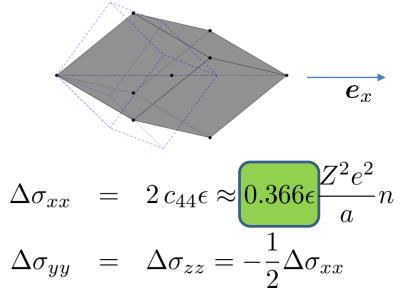
Small deformations: Monocrystal

Fuchs (1936) (neglecting screening):

Volume conserving tension:

$$u_{xx} = \epsilon + \frac{3}{2}\epsilon^2$$
$$u_{yy} = u_{zz} = -\frac{1}{2}\epsilon$$

Along cube diagonal [111]

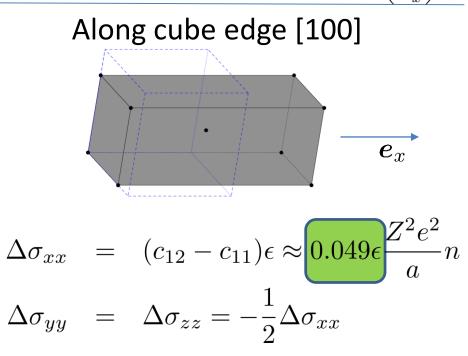


Very strong anisotropy

$$L_{x} = L_{y} = L_{z} = V^{1/3}$$

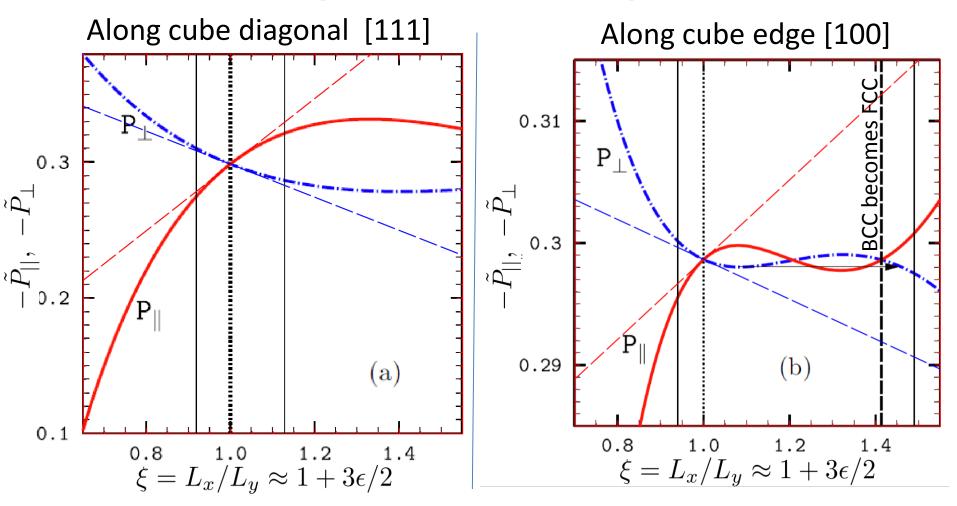
$$L_{z}$$

$$L_$$



Large deformations: Monocrystall

[Baiko&Kozhberov 2017]



Stress is strongly nonlinear (at certain directions) Breaking stress is anisotropic (note the difference in scales)

Small deformations: Polycrystal

Polycrystalline matter, made of large number of single crystals, should be isotropic.

$$\Delta \sigma_{ij} = 2\mu_{\text{eff}} \left(u_{ik} - \frac{1}{3} \delta_{ik} u_{ll} \right) + K u_{ll} \delta_{ik}$$

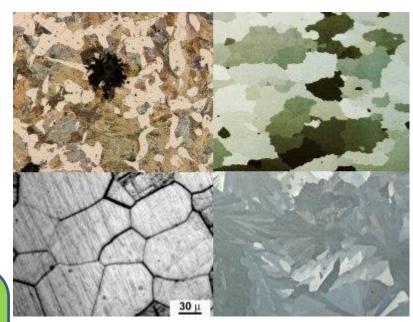
Shear modulus

Compression modulus

Landau&Lifshitz, vol. 7: "The relation between the elastic properties of the whole crystal and those of components crystallites depends on actual form of the latter and the amount of correlation of their mutual orientation."

Ogata&Ichimaru (1990): (averaging velocity of shear waves)

Kobyakov&Pethick (2015): (self consistent theory by Eshelby 1961)



[From Wikipedia]

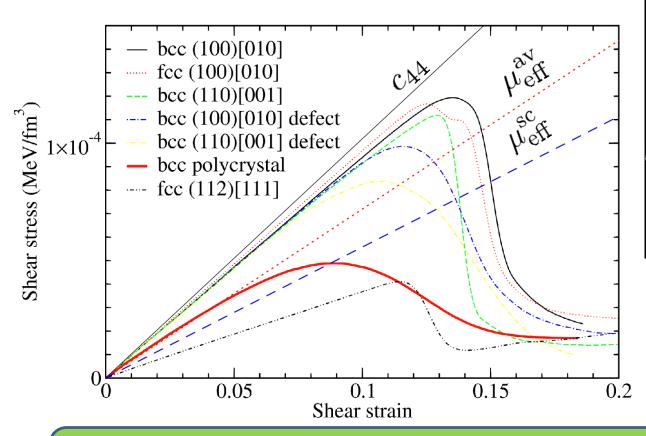
28%

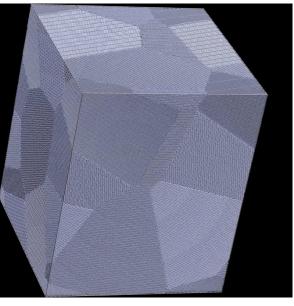
Herenc

$$\mu_{\text{eff}}^{\text{av}} = 0.120 \frac{Z^2 e^2}{a} n$$
$$\mu_{\text{eff}}^{\text{sc}} = 0.093 \frac{Z^2 e^2}{a} n$$

Small deformations: Polycrystal

Numerical experiment: MD simulations by Horowitz&Kadau (2009) for shear deformations



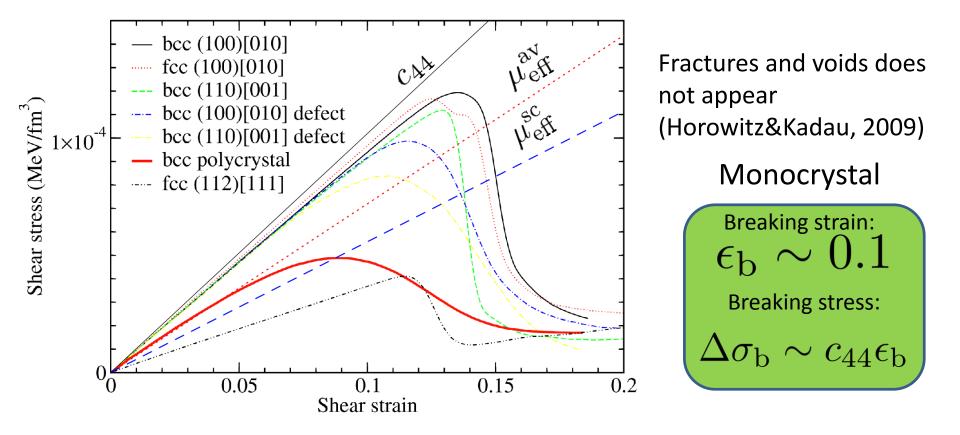


Movie and original figure from Horowitz&Kadau (2009)

Shear stress is almost linear for monocrystal Effective shear modulus is rather uncertain

Breaking strain&stress

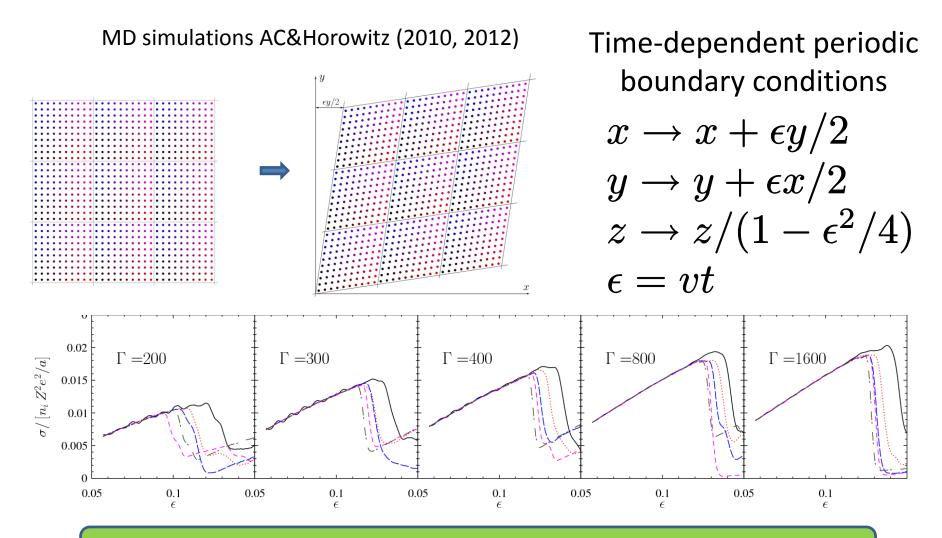
Numerical experiment: MD simulations by Horowitz&Kadau (2009) [shear deformations]



Polycrystal: breaking strain and stress are moderately reduced

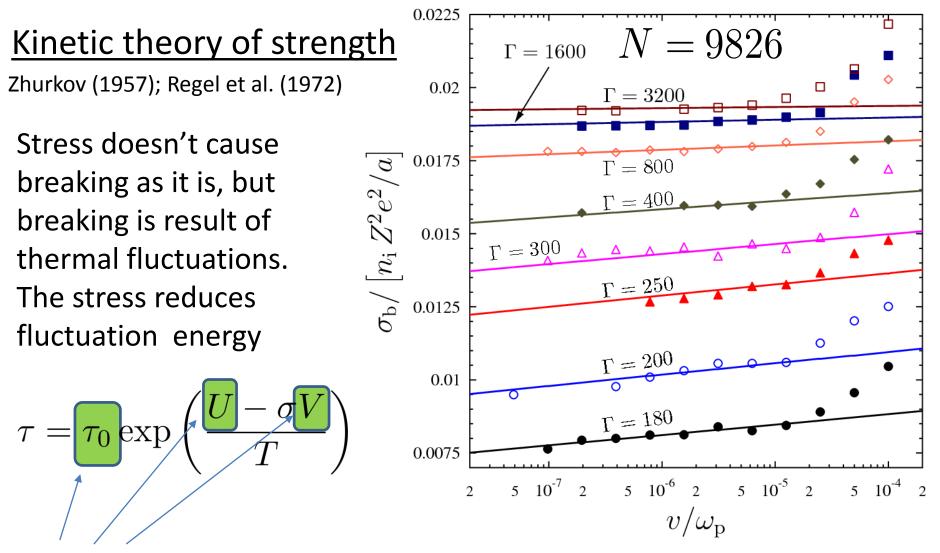
Durability of the crust

How long crust can support the stress?



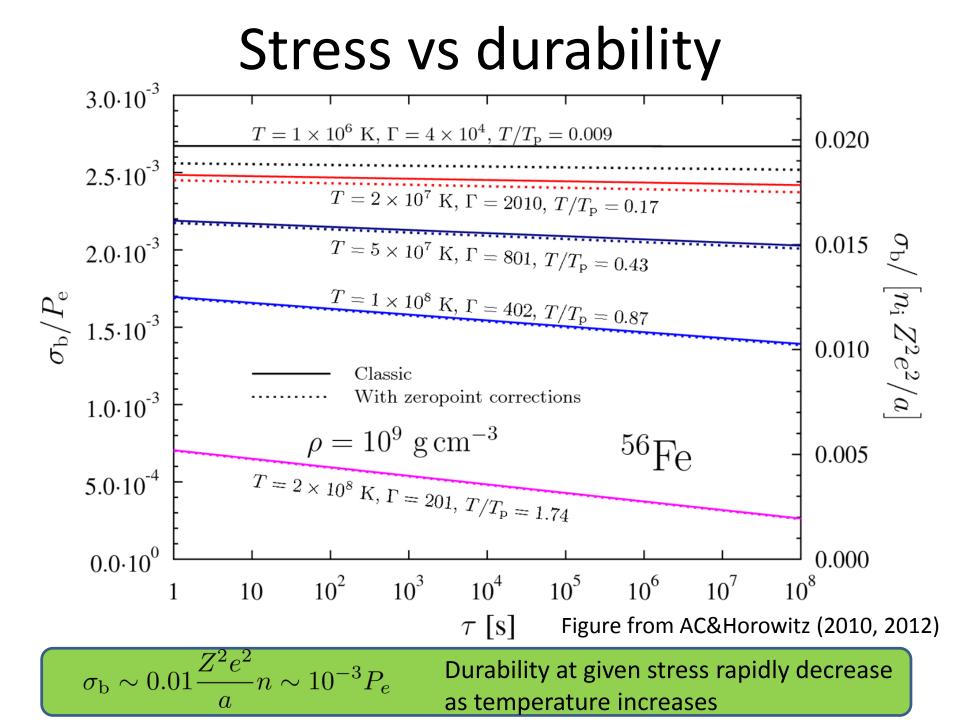
Breaking stress depends on the temperature and strain rate

Durability of the crust



Parameters to be fited to [numerical] experiment

Figure from AC&Horowitz (2010, 2012)



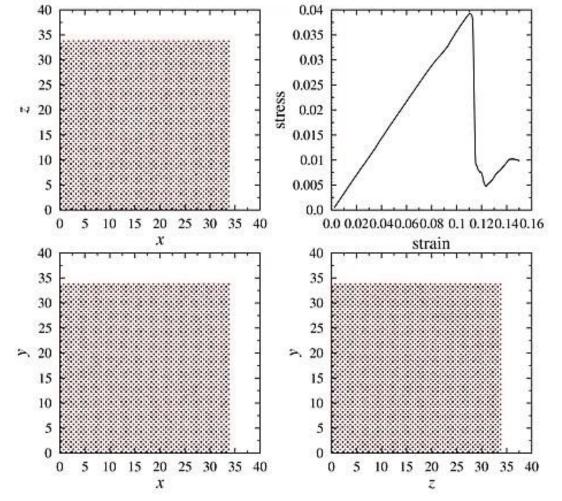
How does the crust break?

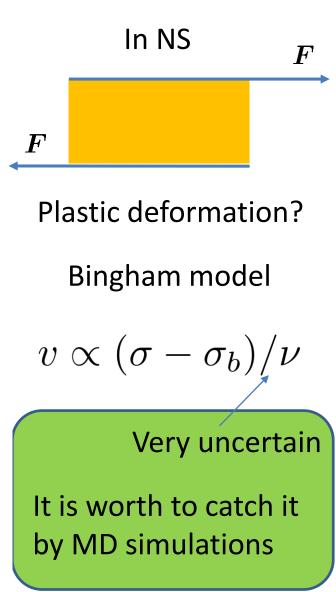
In MD simulations

Simulation: md2q_G800_e0_0.15_t48000.traj

 $\Gamma=\!800,\,N=\!78608,\,Z=\!100,\,\lambda_{\rm c}/a=\!1.17,\,\epsilon=\!0.000780414,\,v/\omega_{\rm p}=\!3.125\text{e-}06$

Data from:md_G800_e0.00000_0.00078_t249.833.out Str_G800_e0_0.15_t48000.out





Summary



- 1. The elastic properties of crystal monocrystals are known. They are anisitropic and nonlinear. Elastic properties of polycrystalline matter are rather uncertain (~28% difference).
- 2. Uniform compression/expansion does not lead to breaking
- 3. MD simulations (Horowitz&Kadau, AC&Horowitz, Hoffman&Heyl) suggest that the crust breaks at shear strain ~0.1, corresponding to stress $\Delta \sigma_{\rm b} \sim 0.01 \frac{Z^2 e^2}{a} n \lesssim 10^{-3} P_e$
- 4. Durability of the crust is a strong function of the stress and decreases with increase of temperature. It can lead to thermoplastic instability (Levin&Beloborodov 2014)
- 5. The real durability can hardly exceed extrapolation of MD simulations [AC&Horowitz 2010,2012], but I would not be strongly surprised, if in fact it is lower (weakness along edge?)

Short answers to the Bob's question